Lecture 3 – Earth surface temperature

ALBEDO

- Albedo: reflection coefficient can range from 0 (no reflection) to 1 (100% reflection).
- On average earth has an albedo of 0.3 in the visible range.
- A (visible) = 0.3: 1/3rd from earth's surface, 2/3rd from atmosphere.
- A (IR) ~ 0

1

2

























Z	eroth Ord	ler Model		
Not a bac Λ _{max} x T _o Ω ά 1/d²	l model ! = constant			
Earth Venus Mars Jupiter	T _o calc 255 227 216 98	T _o meas 255 230 220 130	Tsurf 290 750	
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FEEDBACK	
$\begin{array}{l} 2 \ x \ CO_2 \ evel = 550 \ ppm; \ [275 \ ppm \ x \ 2 = 550 \ ppm] \\ \Delta Ts = 1.4 \ to \ 5.8^{\circ}C \ (this \ big \ range \ because \ of \ feedback) \\ Feedback \\ Output = \ lnput \ + \ g^1! \ + \ g^2! \ + \ g^3! \ \dots \ = 1/(1-g) \\ If \ g > 1 \ instability \\ If \ g < 0, \ 0 < 1. \qquad If \ g > 0, \) > 1. \\ Major \ +ve \ feedback \ impacts: \\ Water \ vapor \ (major \ GHG). \ Warmer \ it \ gets \ more \ water \ vapor \ lce-albedo. \ Warmer, \ less \ ice, \ decrease \ albedo. \\ CO2 \ from \ oceans. \ Higher \ temperature \ more \ outgassing. \\ Clouds \ not \ understood \ and \ reason \ for \ large \ uncertainty. \end{array}$	
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EARTH'S SURFACE TEMPERATURE				
Cor	rective factors we need to put in are:			
1.	Vertical upliftment of convective heat (Fc) – loaded fully into the lower zone			
2.	Latent heat (Fe) – 50% lower and 50% upper region.			
3.	Sunlight absorbed in air (Fs) – 70% in upper atmosphere and 30% in lower.			
4.	Some IR passes through atmosphere (Fw).			
	These factors reduce the surface temperature since they either reduce heat input or increase its rate of removal. $\Omega/4 = a \Omega/4 + \sigma T_o^4 + Fw$ (5)			
	2* σT _o ⁴ = σT ₁ ⁴ + 0.5Fe + 0.7Fs(6)			
	2* σT ₁ ⁴ = σT _o ⁴ + σT _s ⁴ – Fw + Fc + 0.5Fe + 0.3Fs(7)			
Sho	w for a two layer model that:			
	σT _s ⁴ = 3(1-a) Ω/4 – (Fc + 1.5Fe + 1.7Fs + 2Fw)(8)			
	Calculate the adjusted T_{12} , T_{13} , and T_{2} using the numbers from Figure "Trapping of heat by GHG".			
	Answer: Ts = 287K			
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THE CO₂ "GREENHOUSE" EFFECT

Direct effect of an increase in atmospheric CO₂ and other GHG is increase in amount of IR absorbing materials. In our model this results in the increase in n – the number of IR – absorbing boxes. Eqn 8 can be rewritten as: $\sigma T_s^4 = (n+1)(1-a) \Omega/4 - (Fc + 1.5Fe + 1.7Fs + nFw)(9)$ Using the numerical values from "Trapping of heat by GHGs", Ts = $1/\sigma [(n+1) (1-a) \Omega/4 - 283.2 + n(20)]^{1/4}$(10) To find the impact of a small change of n on Ts, **dTs/dn** = (1-a) $\Omega/4\sigma - 40/\sigma / [4\{(n+1)(1-a) \Omega/4 - 283.2 - n(40)\}]^{3/4}$(11) Using eqn (10), this can be rewritten as: **dTs/dn** = Ts/ 4(n+1) * (1 + 263.2/ σT_s^4)(12) Using $\sigma T_s^4 = 397.1 \text{ Wm}^2$ $\Delta Ts/Ts = 0.42 \Delta n/(n+1)$ (13)

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..continued Simple calculation to find the impact of doubling of CO_2 on Δn : Each of the GHGs is effective at absorbing IR at a different part of the spectrum. Certain atmospheric trace gases such as CH4 and O3 can have a particularly high impact on surface temperature. Total umber of moles or air = 1.8×10^{20} . Concentration of CO_2 in atmosphere pre-industrial = 275 ppm(v) Total number of moles of CO₂ = 49.5 x 10^15 ((275/10^6)*1.8*10^20) Total number of moles of $H_2O = 722 \times 10^{15} ((1.3 \times 10^{19}))$ Ratio of CO₂ to (H_2O+CO_2) , r = 0.064 Since a mole of any substances contains Avogadro's number of molecules, r is also the ratio of molecules. An individual molecule of CO₂ is only 0.25 times as effective as a molecules of H₂O in terms of absorbing IR. Therefore r = 0.25 times the earlier value. $\Delta n = r x \%$ increase in CO₂ x n Δn = 0.0169 x 1 x n = 0.0338 Substituting in Eq 13 $\Delta Ts = 0.42/3 \times 0.0338 = 0.0047Ts$ Using the value Ts = 290K, ΔTs = 1.4K Lecture 3 26















- By changing the incoming solar radiation (e.g., by changes in Earth's orbit or in the Sun itself)
- By changing the fraction of solar radiation that is reflected (called albedo) (e.g., by changes in cloud cover, atmospheric particles or vegetation)
- By altering the longwave radiation from Earth back towards space (e.g., by changing greenhouse gas concentrations)

Climate, in turn, responds directly to such changes, as well as indirectly, through a variety of feedback mechanisms.

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Factors affecting temperature

- Clouds:
- Fine particles (aerosols) in the atmosphere:
 - Atmospheric factors shown in the image below include natural factors such as clouds, volcanic eruptions, natural biomass (forest) burning, and dust from storms. In addition, human-induced factors such as biomass burning (forest and agricultural fires) and sulfate aerosols from burning coal contribute tiny particles that contribute to cooling.
 - Volcanic eruptions of this magnitude can impact global climate, reducing the amount of solar radiation reaching the Earth's surface, lowering temperatures in the troposphere, and changing atmospheric circulation patterns. The extent to which this occurs is an uncertainty.
 - Water vapor is a greenhouse gas, but at the same time, the upper white surface of clouds reflects solar radiation back into space. Albedo—reflections of solar radiation from surfaces on the Earth—creates difficulties in exact calculations. If, for example, the polar icecap melts, the albedo will be significantly reduced. Open water absorbs heat, while white ice and snow reflect it.

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Ocean (Te ocean conveyor belt):













